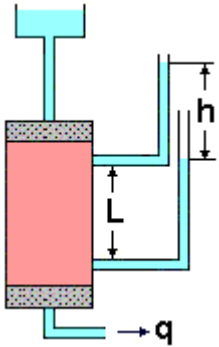


LECTURE 11

Laboratory Measurement of Permeability:

Constant Head Flow

Constant head permeameter is recommended for coarse-grained soils only since for such soils, flow rate is measurable with adequate precision. As water flows through a sample of cross-section area **A**, steady total head drop **h** is measured across length **L**.

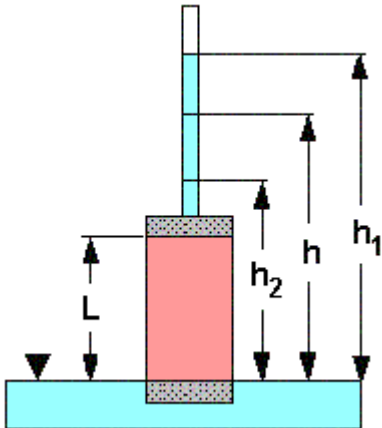


Permeability **k** is obtained from:

$$k = \frac{qL}{Ah}$$

Falling Head Flow:

Falling head permeameter is recommended for fine-grained soils.



Total head **h** in standpipe of area **a** is allowed to fall. Hydraulic gradient varies with time. Heads **h₁** and **h₂** are measured at times **t₁** and **t₂**. At any time **t**, flow through the soil sample of cross-sectional area **A** is

$$q = k \cdot h \cdot \frac{A}{L} \text{ ----- (1)}$$

Flow in unit time through the standpipe of cross-sectional area **a** is

$$= a \times \left(-\frac{dh}{dt} \right) \text{-----} (2)$$

Equating (1) and (2) ,

$$-a \cdot \frac{dh}{dt} = k \cdot h \frac{A}{L}$$

$$\text{or } -\frac{dh}{h} = \left(\frac{kA}{La} \right) dt$$

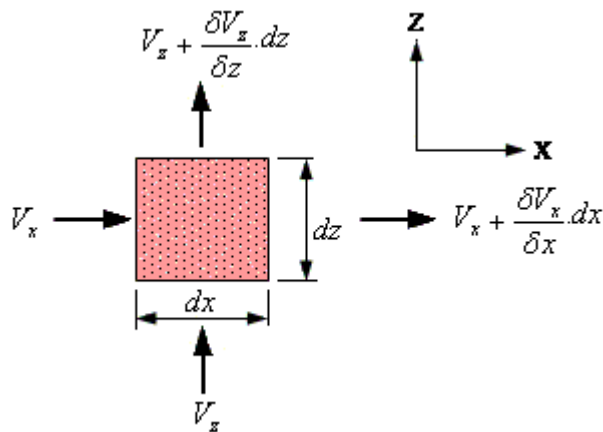
Integrating between the limits,

$$\log_e \left(\frac{h_1}{h_2} \right) = \frac{k \cdot A}{L \cdot a} (t_2 - t_1)$$

$$k = \frac{L \cdot a \cdot \log_e \left(\frac{h_1}{h_2} \right)}{A(t_2 - t_1)}$$

$$= \frac{2.3 L \cdot a \cdot \log_{10} \left(\frac{h_1}{h_2} \right)}{A(t_2 - t_1)}$$

Seepage in Soils:



A rectangular soil element is shown with dimensions **dx** and **dz** in the plane, and thickness **dy** perpendicular to this plane. Consider planar flow into the rectangular soil element.

In the **x-direction**, the net amount of the water entering and leaving the element is

$$\frac{\delta V_x}{\delta x} . dx . dy . dz$$

Similarly in the **z-direction**, the difference between the water inflow and outflow is

$$\frac{\delta V_z}{\delta z} . dz . dx . dy$$

For a two-dimensional steady flow of pore water, any imbalance in flows into and out of an element in the z-direction must be compensated by a corresponding opposite imbalance in the x-direction. Combining the above, and dividing by $dx . dy . dz$, the **continuity equation** is expressed as

$$\frac{\delta V_x}{\delta x} + \frac{\delta V_z}{\delta z} = 0$$

From Darcy's law, $V_x = k_x . \frac{\delta h}{\delta x}$, $V_z = k_z . \frac{\delta h}{\delta z}$, where **h** is the head causing flow.

When the continuity equation is combined with Darcy's law, the equation for flow is expressed as:

$$k_x . \frac{\delta^2 h}{\delta x^2} + k_z . \frac{\delta^2 h}{\delta z^2} = 0$$

For an isotropic material in which the permeability is the same in all directions (i.e. $k_x = k_z$), the **flow equation** is

$$\frac{\delta^2 h}{\delta x^2} + \frac{\delta^2 h}{\delta z^2} = 0$$

This is the **Laplace equation** governing two-dimensional steady state flow. It can be solved *graphically, analytically, numerically, or analogically*.

For the more general situation involving *three-dimensional* steady flow, Laplace equation becomes:

$$\frac{\delta^2 h}{\delta x^2} + \frac{\delta^2 h}{\delta y^2} + \frac{\delta^2 h}{\delta z^2} = 0$$